

Prediction of Flood Area Based on the Occurrence of Rainfall Intensity

Agus Suharyanto

Abstract— Flood is one of natural hazard often occurs in many areas. During raining season flood is occur in many areas in Indonesia. There are many reasons of flood occurring, among them are high rainfall intensity, degradation of watershed condition, illegal logging, and increasing of percentage of impervious area especially in urban district. Usually, the flood is occur because of the flood water level is higher than embankment level. The area have elevation lower than flood water level will be inundated by the flood. If the watershed condition is not changed, the water level is only depend on the rainfall intensity. In this research, the area of inundation due to the rainfall intensity data will be investigated. The rainfall intensity was analyzed based on certain return period. The water level in channel was analyzed using HEC-RAS software. The water level is plotted in the topographic map to predict the inundation area. By selecting the rainfall intensity data which is same with rainfall intensity when the flood was occurred as input data, the inundation area was investigated. The investigated result of flood inundation map was compared with flood inundation map generated from observation data during the flood. It was can be confirmed that the investigated and surveyed flood maps are similar. Therefore, it can be concluded that analyzing rainfall data with certain return period is powerful to predict the inundation of flood.

Keywords — About four key words or phrases in alphabetical order, separated by commas.

1. INTRODUCTION

During raining seasons, flood is occurring in many places in Indonesia. The main reason is the surface runoff discharge is bigger than channel capacity. Due to the change of physical condition of watershed and climate change, every year the runoff discharge is increased. On the other way, every year the channel capacity is decrease due to the sedimentation phenomena. To predict the flood inundation area, it is necessary to select the rainfall intensity data same with the rainfall intensity which is causing the flood in an area. Usually, the rainfall intensity data used as input data in the runoff simulation was selected from rainfall intensity with certain return period. The rainfall intensity data used as input data to the runoff simulation have same value with recorded rainfall intensity when the flood was occurred. This data is selected from rainfall intensity with certain return period. To predict the water level the HEC-RAS software is used. To run the HEC-RAS, beside of rainfall data the channel cross section, land use, roughness coefficient, watershed area, and channel long section data are need. Channel cross section was collected from topographic survey. The land use and roughness coefficient data were generated from satellite remote sensing data. The watershed area was analyzed from topographic map. For the rainfall data, 10 years time series data were used. Daily rainfall intensity data was used and daily maximum data is selected to analyze the rainfall intensity for many return period years. The

analyzed result of rainfall intensity in many return period it was in daily rainfall intensity. Due to the analyzing process data requirement, it is necessary to change the daily rainfall intensity to hourly rainfall intensity. The Mononobe formulas was used to change the daily rainfall to hourly rainfall data [1].

The history of flood occurrences of the research area was recorded by the relating agency. Two biggest floods were selected as samples in the research analysis. By looking the dates of the flood occurrences, the rainfall data at the same dates were investigated. The recorded rainfall intensity data at the day when the flood was occurrences was compared with rainfall intensity were analyzed at the many return periods. The years of the return period of recorded rainfall was decided when the values of both rainfall intensity are same. By using the selected rainfall intensity data in certain return period, the water level in the channel were simulated using HEC-RAS software. The others data need in the analysis such as land use, watershed characteristics, and channel characteristics were generated from the raw data collected with the same year. From the simulation result, the water levels at cross and long sections of channel were obtained. The inundation area were plotted at the topographic map due to the elevation of simulated water level result. The boundary of inundation area was represented by the contour line have elevation with the simulated water level elevation. The area of inundation from the simulation was compared with the area of inundation drew from the field survey. The agreement of the comparison can be confirmed well. Therefore, the rainfall data analyzed by certain return period can be used to predict the flood inundation in an area.

Agus Suharyanto is with the Civil Engineering Department, Faculty of Engineering, Brawijaya University Jl. MT. Haryono 167, Malang 65145, Indonesia. Phone: +62-341-580120; Fax.: +62-341-551430; E-mail: agus.s@ub.ac.id.

2. RESEARCH MATERIALS AND METHOD

The research area is Sampean watershed, which is located in east Java province Indonesia. The watershed occupied two regencies, namely Bondowoso regency and Situbondo regency. Geographically, Sampean watershed located between 7.65° - 8.125° South Latitude and 113.65° - 114.125° East Longitude. The area of Sampean watershed is 1,239.77 km². To analyze the rainfall intensity in many return periods, surface runoff discharge, and water level in channel many data and materials are needed. The rainfall intensity in many return periods is analyzed using daily rainfall data in 10 years time series from 27 rainfall gauges. The rainfall stations distribution map is shows in Figure 1. The physical parameters of watershed was generated from topographic map using GIS technology [2]- [3]. The land use data was classified from SPOT remotely sensed data using unsupervised classification method [4].



Fig. 1. Rainfall Distribution Map in Sampen Watershed

The 15 sheets of topographic maps in scale 1:25,000 were used in this research. The SPOT data with path and raw 296/366 and 297/366 scanned in September 2014 were used in this research.

To know the surface runoff discharge, Nakayasu Synthetic Unit Hydrograph (SUH) was used in this research. The Nakayasu SUH formulas are shown from equation 1 to 8 [5]. Typical of Nakayasu hydrograph is shown in Figure 2. The equation 1 is used to analyze the peak discharge. The discharge on rising limb area is analyzed using equation 2 and discharge on decreasing limb is calculated using equation 3-5. The next data need in this research is channel geometry and collected from field survey. By using peak runoff discharge and channel geometry as input data, the water level in channel is analyzed using hydraulic formula [6]. The HEC-RAS software was applied in water level profile analysis in channel [7].



Fig. 2. Typical Hydrograph of Nakayasu SUH.

$$Qp = \frac{C.A.Ro}{3.6(0.3Tp.T0.3)}$$
(1)
$$Qa = Qp \left(\frac{t}{Tp}\right)^{2.4}$$
(2)

where:

Qa = rising limb discharge ($m^3/se.$)

t = time (sec. or hour)

 $Qp = peak discharge (m^3/sec.)$

C = surface runoff coefficient

A = watershed area (km^2)

 $R_o = rainfall unit (mm)$

 T_p = time to peak (hour)

 $T_{0,3}$ = time need in decreasing from peak discharge to 30% of peak discharge.

The part of decreasing limb can be calculated using equation (3-5).

$$Qd > 0.3Qp \qquad Qd = Qp.0.3 \tag{3}$$

$$0.3Qp \ge Qd > 0.3^2 Qp \qquad Qd = Qp.0.3^{\frac{t-Tp+0.5T0.3}{1.5T0.3}}$$
(4)

$$Qd \le 0.3^2 Qp$$
 $Qd = Qp.0.3 \frac{t - Tp + 1.5.T0.3}{2.T0.3}$ (5)

$$Tp = tg + 0.8 tr$$
(6)

If
$$L < 15$$
 km. tg = 0.21 $L^{0,7}$ (7)

If
$$L > 15 \text{ km}$$
 tg = 0,4 + 0,058 L (8)

where:

L = length of channel (km) tg = time concentration tr = 0.5 tg to tg (hour)

 $T_{0,3} = \alpha \text{ tg (hour)}$

tr = flood unit time (hour)

 α = hydrograph parameter

$$\alpha = \frac{0.47(A \cdot L)^{0.25}}{tg}$$
(9)

The rainfall intensity usually recorded in daily data. Due to the data requirement to analyze the runoff discharge, it is necessary to convert the daily to hourly rainfall data. The Mononobe formula is used in this analysis. The formula can be shown in equation 10 [8]-[9].

$$I = \frac{I_{24}}{24} \left[\frac{24}{t_c} \right]^{\frac{2}{3}}$$
(10)

where:

I = rainfall intensity (mm/hour) R_{24} = daily rainfall intensity (mm/day) tc = rainfall duration

As shown in Figure 1, there are 27 daily rainfall data collected from 27 stations were used in this research. One maximum daily rainfall data for each year was selected to use in the analysis. Ten years daily rainfall data was collected in this research. Consequently, there are 27 maximum daily rainfall data for each year. To find the average rainfall intensity for each year in the Sampean watershed, the Thiessen polygon method is used [10]. The Thiessen polygon figure can be shown in Figure 1. Finally, 10 average daily rainfall data in Sampean watershed for each year can be found. From this data the rainfall with certain return period was calculated using Log Pearson Type III.

To convert the daily rainfall data with certain return period to hourly rainfall data, the Mononobe formula was used. This data will be used as one main data to find the surface runoff discharge in the research area. Beside it, the surface runoff coefficient and length of river data are needed in this analysis. The surface runoff coefficient was estimated using land cover map generated from satellite remote sensing data. The length of river was calculated based on topographic map. The total discharge flowing through the river are coming from surface runoff discharge and discharge from domestic wasted water. The discharge of domestic wasted water was analyzed due to the population density. The total discharge for each rainfall data with certain return period will be used as one of main input data to calculate the water level in river. The HEC-RAS software is used in this analysis [11].

The data need in the water level simulation using HEC-RAS are channel cross sectional, channel roughness coefficient, longitudinal slope, and runoff discharge with many rainfall intensity return period. After analyzing process, the water level profile on the river can be found. The water level found from the analysis result is superimposed with the cross sectional data for each cross sectional point with 100 m interval. From the superimposed results it can be found the point which are the water level are higher than bank level. If the water level higher than bank level, it can be concluded that the flood is occurred. By plotting the water level from cross section point which are the flood occurred to the topographic map, the flood area can be predicted. From the simulations results with rainfall intensity in many return period as one of main input data, the certain rainfall return period data which is causing the flood occurrence can be found. The principle of the whole research method can be summarized in flow

diagram as shown in Figure 3.

3. RESULTS AND DISCUSSION

The rainfall data was collected from 2005 to 2014. The maximum daily rainfall data during one year for each rainfall station was selected. There 10 maximum daily rainfall data (from 2005 - 2014) for each station. From the 27 rainfall stations, the 270 maximum daily rainfall data was found. Tabel 1 is shown the 270 maximum daily rainfall data. After consistency test done, the average rainfall data was analyzed using Thiessen polygon method. From the area measurement of Thiessen polygon analyzing result, the influence area of each rainfall station was found. By using the Thiessen polygon formula the average maximum rainfall data for each year in Sampean watershed was found. The average maximum rainfall data is shown in Table 2. From this data the appropriate distribution type was analyzed base on statistical method [12]. The Log Pearson type III is the most appropriate distribution type for the rainfall data on Sampean watershed. Consequently, the rainfall intensity with many return period were analyzed using Log Person type III formula. The analyzed result of daily rainfall intensity in 2, 5,10, 25, 50, 100, and 200 years return period is shown in Table 3.

Table 1. Maximum Daily Rainfall Data

No.	Station Name	2005	2006	2007	2008	2009	2010	2011	2012	2013	2014
1	Pengairan	116	118	108	126	65	96	120	90	100	125
2	Ancar	92	86	175	108	68	92	95	96	185	110
3	Klabang	200	116	106	189	80	125	290	125	110	190
4	Selolembu	61	90	70	166	85	80	72	100	75	174
5	Wonosari II	180	91	93	195	98	93	185	92	95	182
6	Grujugan	111	89	107	195	98	60	121	90	110	175
7	Maesan	103	101	81	137	83	104	93	100	82	130
8	Sukokerto	75	85	90	97	96	77	85	88	85	90
9	WonosariI	81	80	85	117	94	68	80	85	80	110
10	Wonosroyo	87	90	58	130	75	72	85	95	60	140
11	Kejayan	51	95	87	87	75	72	61	100	85	80
12	Kasemek	43	95	82	89	82	52	44	93	80	85
13	Sbr. Gading	83	91	66	92	98	144	85	92	64	95
14	Pakisan	60	68	86	119	56	78	62	70	85	120
15	Tlogosari	68	67	83	98	79	70	67	75	88	100
16	Pinang Pahit	55	93	94	149	82	83	56	90	90	138
17	Maskuning Wtn.	49	56	78	92	60	50	48	60	86	96
18	Prajekan	56	97	74	263	150	89	65	90	88	82
19	Jeru	60	75	61	135	45	60	70	80	72	85
20	Pandan	80	85	60	245	62	80	80	88	65	125
21	Glendengan	68	118	86	141	65	96	78	120	78	110
22	Talep	65	92	78	275	85	89	55	95	82	125
23	TAAL	60	76	62	125	46	65	55	85	65	90
24	Pringduri	67	93	87	134	70	95	66	95	80	85
25	Bluncong	62	117	87	270	65	95	65	125	95	58
26	Kolpoh	61	91	76	175	78	86	60	95	80	115
27	Sulinh With	04	80	52	205	64	100	00	00	55	05

The hourly rainfall intensity is required to input data for analyzing runoff discharge using Nakayasu SUH. Therefore, it is necessary to convert the daily rainfall data to hourly rainfall data. The Mononobe formula was used to convert daily to hourly rainfall intensity data. Base on the recorded rainfall data, the average of rainfall duration in Sampean watershed is 6 hours. Therefore, *tc* (rainfall duration) on Mononobe formula is 6. The hourly rainfall data due to the Monobe formula can be shown in Table 3.



Fig. 3. Flow Diagram of Research Method

Table 2. Average Maximum Rainfall Intensity

Year	I (mm)	Year	I (mm)
2005	93.743	2010	140,237
2006	92.586	2011	86,245
2007	82.513	2012	125,795
2008	149.276	2013	102,125
2009	95.75	2014	95,713

Table 3. Rainfall intensity in Many Return Period

Return Period (year)	I (mm/day)
2	95.215
5	111.634
10	135.788
25	155.325
50	180.121
100	197.315
200	225.855

The next data need to calculate the runoff discharge using Nakayasu SUH are watershed area, length of river and runoff coefficient. The watershed boundary was analyzed due to Digital Elevation Model (DEM) generated from digital topographic map with gridding size 50 x 50 m [13]. The GIS was used to find the watershed boundary. From the measurement result the area of Sampean watershed is 1,239.77 km². The length of river (Sampean river) was measured from digital topographic map in scale 1:25,000. The length of Sampean river it was found equal to 75.5 km. The runoff coefficient was analyzed due to the land cover map. SPOT imagery was used to generate the land cover map [14]. The Isoclass unsupervised classification method was used to classify the land cover map. The land cover were classified to the 8 categories i.e. water, forest, agriculture land, shrub and brush, open land, high density, medium density, and low density. The runoff coefficient for each land cover category was estimated from runoff coefficient table as shown in Table 4. The average of runoff coefficient for Sampean watershed was calculated using weight average method. The calculation result shows the average runoff coefficient for Sampean watershed is 0.565. Based on watershed area, length of river, and runoff coefficient as input data to Nakayasu SUH the unit hydrograph was calculated with 1 mm rainfall unit. The discharge of unit hydrograph can be found as shown in Table 5. Due to the discharge of unit hydrograph, the runoff discharge with rainfall intensity in many return periods was calculated. The peak discharge for many rainfall intensity in many return periods can be shown in Table 6. The hydrograph is shown in Figure 4.

In Table 4 is shown that the peak discharge will be occurred seventh hours after raining was begun. The time base of hydrograph is 25 hours. It means the peak discharge will occur not so fast. This phenomena indicated that the Sampean watershed still have good environment condition. It is agree with the existing condition, where the land cover condition are dominated by forest and agriculture land. Figure 4 shows that the rising limb is faster than decreasing limb. This data indicated that the characteristic of Sampean watershed are high land slope, the roughness coefficient is little bit high, and manmade land cover is growing. Base on this hydrograph, on the future it is necessary to minimize the peak discharge and increasing the peak time. Therefore, better management land cover management is become one of most important issue on the future.

Table 4. Runoff Coefficient

No.	Land Cover Category	С
1	Water	0.1
2	Forest	0.2
3	Agricultural land	0.45
4	Shrub and Brush	0.5
5	Open Land	0.3
6	High Density	0.8
7	Medium Density	0.7
8	Low Density	0.55

No. Time (hour)		Unit Discharge (m ³ /sec.)			
1	0.00	0.000			
2	1.00	0.250			
3	2.00	1.183			
4	3.00	3.175			
5	4.00	6.535			
6	5.00	10.485			
7	6.00	17.136			
8	7.00	24.517			
9	8.00	18.885			
10	9.00	17.364			
11	10.00	15.973			
12	11.00	14.655			
13	12.00	13.515			
14	13.00	12.428			
15	14.00	11.459			
16	15.00	10.692			
17	16.00	9.675			
18	17.00	8.882			
19	18.00	8.152			
20	19.00	7.475			
21	20.00	7.102			
22	21.00	6.225			
23	22.00	5.815			
24	23.00	5.360			
25	24.00	4.866			

Table 5. Runoff Coefficient



Fig. 4. Nakayasu Hydrograph.

The runoff discharge is flowing through the river and indicated by river water level. It is assumed that all of the runoff discharge is flowing to the river. By inputting the peak discharge, channel geometry, river roughness coefficient, and others need data the water level in channel was analyzed using HEC-RAS. The river cross section was surveyed with 100 m interval and was numbered from 1 to 77 starting from upstream to downstream. The layout of cross section point for Sampean river is shown in Figure 5. After water level profile simulation using HEC-RAS, the water level profile for each cross section can be found. From each water level profile in each cross section figure, the water level and river bank positions were investigated. If the water level position is higher than river bank level, the flood is occurred in that cross section point. For examples, Figure 6 and Figure 7 are shown the water level in cross sections point 10 and segment 49 with rainfall data 100 year return period. From this figures it can be seen that the water level are over of the river bank elevation. Therefore, from these cross sections the flood will be occurred. From the analyzed result, there are many cross section points with situation that water level are higher than river banks elevation. From the water level higher than river bank elevation, the inundation area was drawn in the topographic map. Contour elevation is as guide to draw the inundation area.

Table 6. Flood Discharge in Many Return Period

Return Period	Flood Disch. (m ³ /sec.)	Return Period	Flood Disch. (m ³ /sec.)
2	865,105	50	1.519,850
5	1.060,405	100	1.855,575
10	1.205,218	200	2.082,405
25	1.430,176		



Fig. 5. Sampean River Cross Section Points.



As mentioned above, the water level simulation was done by inputting the rainfall intensity in many return periods. The rainfall intensity with 2, 5, 10, 27, 50, 100, an 200 years return period were used as input data in water level simulation using HEC-RAS. By interpreting the position of water level from simulation result and river bank level, the flood occurrences were investigated. From the investigation result it can be concluded that in Sampean watershed the flood will be occurred if the rainfall intensity is bigger or equal than rainfall intensity with 50 year return period. This conclusion are agree with the flood were occurred in 2008, 2010, and 2012. In that moments, the recorded daily rainfall intensity were 180 mm, 200 mm, and 195 mm respectively. The daily rainfall cause of flood in Sampean watershed on 2008, 2010, and 2012 are higher than rainfall intensity with 50 years return period and nearest with rainfall intensity with 100 year return period as shown in Table 3. The number of cross section which are the water level higher than river bank are 31 cross sections. These points of cross section number are 5,7-12, 53-60, and 61-77, respectively. This simulation result indicated that the downstream area should be wary if in the watershed have rain with intensity more than 180 mm/day or 98 mm/hour.



Fig. 7. Water Level in Cross Section 49.

To draw the flooded area caused by runoff discharge with rainfall intensity input data equal to rainfall intensity in 100 year return period, the water level elevation in each cross section point was plotted to the topographic map. Only in the cross section points which the flood occurred, the flooded area was drew. The flooded area was drew by tracing the contour which are the contour have same elevation with the water level elevation in the river. The map of flooded area and call as Sampean watershed flood map due to the rainfall intensity with 100 years return period can be shown in Figure 7. From that figure it can be seen that the flood was occurred almost in the left side river bank. From the topographic map it is shown that the left side of river bank almost with topographic condition in low land condition. The point to check the validity of Sampean watershed flood map, this map was compared with flood map drew based on the field survey. The field survey was done by water resources local agency when the flood occurred. The comparison was done by overlapping flood maps drawn based on simulation result and drawn due to the field survey. The both of flood maps was drawn with digital format using UTM coordinate system. Symmetric difference overlay technique on GIS program was used in this analysis. The agreement both of maps were shows from the overlapping result.



Fig. 8. Sampean Watershed Flood Area.

4. CONCLUSIONS

From the analyzed results it can be concluded that:

- 1. Rainfall data in many return period is useful to predict the flood inundation.
- 2. Rainfall intensity with 100 year return period it was bringing on the flood in 2008, 2010 and 2012 at Sampean watershed.
- 3. Left side bank of Sampean river should be wary of flood during raining season.
- 4. It is necessary to minimize land cover changed relating with decreasing peak discharge and increasing peak time discharge.

ACKNOWLEDGMENT

The financial of this research was supported by Situbondo local regency. Researcher would like to say many thank for Water Resources agencies of Situbondo and Bondowoso regencies for they are full supported the hydrological data. And many thank for Brawijaya University for supporting the budget to run this research.

REFERENCES

- Chow, V.T., D.R. Maidment, D.R., Mays, L.W. 1988. *Applied Hydrology*, McGraw-Hill Book Co., Singapore.
- [2] Suharyanto, A. Application of Satellite Remote Sensing Data and GIS to Predict the Flood Discharge. In Proceedings of The 9th International Conference on Quality in Research (QiR), Faculty of Engineering, Universitas Indonesia, Jakarta , Indonesia, 6-7 September.
- [3] Suharyanto, A., Suhartanto, E., and Pudyono 2010. The Use of Satellite Remote Sensing data and Geographic Information Systems on Critical Land

Analysis. *Journal of Agrivita*, ISSN: 0126-0537: Faculty of Agriculture, Brawijaya University.

- [4] Aburu S., Golla S. B., 2015. Satellite Image Classification Methods and Techniques: A Review, *International Journal of Computer Applications*, Vol. 119., No.8.
- [5] Safarina A. B., Salim H. T., Hadihardaja I. K., Syahril M. BK., 2011. Clusterization of Synthetic Unit Hydrograph Method Based on Watershed Characteristics, *International Journal of Civil & Environmental Engineering* (IJCEE-IJENS), Vol. 11., No.6.
- [6] Plate, E.J., 2002. Flood Risk and Flood Management, *Journal of Hydrology*.
- [7] Mays L. W., 2005. *Water Resources Engineering*, John Wiley and Sons, USA.
- [8] Sosrodarsono, S. and Takeda, K. 2003. *Hydrology for Water Resources*, Pradnya Paramita Publisher, Jakarta, Indonesia.
- [9] Pardian, M. and Kumanan, C.J., 2014. Soil Erosion Vulnerability Mapping Using Remote Sensing Based MMF Rule in Part of Coimbatore and Tiruppur-Districts Tamil-Nadu India, *International Journal of Geomatics and Geosciences*, Vol. 5., No.1., ISSN: 0976-4380.
- [10] McCuen, R. H., 1988. Hydrologic Analysis and Design, Prentice-Hall, Inc., Upper Saddle River, New Jersey, USA.
- [11]Brunner, G.W., 2014. Combined 1 D and 2 D Modeling with HEC-RAS, *http://www.hec.usace. army.mil/misc/files/ras.*
- [12] Chow, V.T., 1988. *Applied Hydrology*, McGraw-Hill Series in Water Resources and Environmental Engineering, Boston, USA.
- [13] Singh, V.P. and Florentino, M., 1996. *Geographical Information Systems in Hydrology*, Kluwer Academic Publishers, London, UK.
- [14] Starck, J.L., Murtagh F., and Bijaoui A., 1998. *Image Processing and Data Analysis*, The Multiscale Approach, Cambridge University Press, UK.